



## A REVIEW OF RADIOMETRIC ANALYSIS ON SOIL EROSION AND DEPOSITION STUDIES IN AFRICA

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**Abstract:** Soil erosion is one of the main soil degradation phenomena that threaten sustainable use of soil productivity thus affecting food security. In addition, it leads to reservoir storage capacity loss because of sedimentation. This not only affects water quantity but also water quality. Worldwide, annual loss in reservoir storage capacity due to sedimentation is 0.5 to 1%. Similarly, about 27% of land in Africa is largely degraded by erosion. As a result, there is need to minimize soil erosion and deposition through site specific estimation of soil erosion and deposition rates in the reservoirs. To achieve this, Fallout RadioNuclides (FRNs) are some of the methods in use. The most common radionuclides include; <sup>137</sup>Cs, <sup>210</sup>Pb and <sup>7</sup>Be. Only few countries in Africa have exploited these FRNs. In these countries, <sup>137</sup>Cs has been largely exploited but in some regions, it has been reported to be below minimum detection limit. Using <sup>137</sup>Cs and <sup>210</sup>Pb, maximum reference inventory in Africa is found to be 1450 and 2602 Bq/m<sup>2</sup>, respectively. However, there is minimal application of <sup>7</sup>Be within the continent. Also, very little has been done in Africa to assess chronology and sedimentation rates of reservoirs using FRNs measured from sediment cores. In conclusion, a gap still exists on FRNs application in Africa in assessing soil erosion, deposition and reservoir sedimentation.

**Keywords:** soil erosion, deposition, fallout radionuclides, Caesium-137, Lead-210, Beryllium-7.

### 1. INTRODUCTION

Worldwide, agriculture is not only being affected by global warming and climate change but also accelerated soil erosion and the associated soil degradation (Walling, 1998; Marestoni *et al.*, 2009). Accelerated soil erosion is one of the main environmental problem (Zapata and Agudo, 2000; Konz *et al.*, 2010; Othman and Ismail, 2012) that largely affects sustainable use of soil and water resources (Patrocinio and Andrello, 2009). Soil erosion is a major threat to soil productivity and represents the main

mechanism of land degradation worldwide (IAEA, 2014). According to ICOLD (2009), 27% of land in Africa is degraded, out of which 65% of agricultural land is degraded due to erosion and/or chemical and physical damage. In addition, soil erosion is reported by Collins *et al.* (2001) as a serious environmental problem in many African countries. The severe soil erosion and degradation largely affect food security in Sub-Saharan Africa (Dercon *et al.*, 2012). This has been attributed to many developing countries in Africa having experienced expansion of agriculture on to marginal land. To tackle the loss of productive soil, quantification of soil erosion rates is a major requirement (Poręba, 2006; Porto *et al.*, 2014b). Further, soil erosion and deposition have offsite effects related to loss of reservoir storage capacity. In water resources, most reservoirs do not fulfil their func-

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tions due to capacity loss and reduction in water quality resulting from sedimentation (Marestoni *et al.*, 2009; Apostu *et al.*, 2012).

Globally, the loss of reservoir storage capacity has been approximated to 0.5–1% per year (ICOLD, 2009). This greatly affects the reservoir functions like water supply for domestic, irrigation and hydropower generation. On the other hand, water quality is lowered especially when the source of sediments is from agricultural farms or contaminated environment. The eroded sediment acts as a vector for the transportation of agrochemicals from agricultural lands to water bodies (Porto and Walling, 2012). Further, due to climate change and global warming, water erosion is expected to be on the increase (Mabit *et al.*, 2014c). An increase in extreme daily rainfall has been experienced in many catchments and this may lead to severe loss of soil in cultivated land. As a result, there is need to estimate the chronology of sediment production especially in most catchments within developing countries where the sedimentation rates of reservoirs are inadequately documented. Also, soil erosion rate data at different temporal changes is required so as to aid in quantifying potential changes that might occur when different land uses have been adopted and with change in climate (Li *et al.*, 2010; Benmansour *et al.*, 2013).

To reconstruct the history on impacts of anthropogenic activities in water bodies, sediment collected from lakes play a major role. The collected sediments also aids in establishing accurate chronologies of sedimentation and in study of sediment accumulation rates (IAEA, 2003). The use of radionuclides in soil erosion and deposition studies allows assessment of soil erosion and deposition within a catchment to be obtained with minimal field visits (IAEA, 1998; Walling, 1999; Saç and İçhedef, 2015). The Fallout radionuclides (FRNs) is a cost-effective tool that is useful in studying soil redistribution due to erosion within the landscape from plot to basin scale. Also, the information derived from FRNs is useful in complementing information provided by conventional erosion measurements and modelling (IAEA, 2014).

The traditional methods of quantifying soil erosion rate are labour intensive and require long term monitoring programs (Poręba, 2006). Walling and Quine (1992), reported that existing soil erosion monitoring techniques has many limitations in terms of their potential to provide information on longer term soil loss rates over large areas. In addition, in most developing countries especially in Africa, continuous sediment gauging has not been done consistently for longer periods. To tackle these challenges, FRNs have been adopted in many developed countries to assess soil erosion magnitudes (Zapata and Agudo, 2000; Dercon *et al.*, 2012; Porto and Walling, 2012) but little has been done in Africa. Therefore, there is need to explore the use of FRNs in assessing soil erosion rates and patterns of soil redistribution/deposition. The theory of using isotopes in soil erosion analysis is based on their

characteristics. The isotopes reach the soil through precipitation and they are then strongly bound to the soil particles where they can only be moved by erosion (Saç and İçhedef, 2015). The use of radionuclides in determining soil erosion rates, aids in modelling soil redistribution within the catchment or the landscape (Nouira *et al.*, 2003; Porto *et al.*, 2014b). Further, use of the radionuclides provides a fast and robust way of determining spatial soil redistribution (Meusburger *et al.*, 2013). Use of FRNs, provides data that is useful in targeting the areas with high erosion and sediments transfers. The identified areas can then be targeted for land management and conservation practices thus reducing the negative impacts resulting from soil erosion. Successful use of FRNs require suitable and statistically sound sampling procedures (Mabit *et al.*, 2013). This highly reduces the limitations associated with the use of FRNs in soil redistribution studies. Care and proper selection of reference sites is required for proper application of FRNs in assessing soil redistribution. In addition, there is need for more research in various parts of the world on use of FRNs and assessing their limitations as challenged by Parsons and Foster, 2011. This would lead to an improved and highly viable method of assessing erosion and deposition.

In Africa and most developing countries, very few studies have been conducted in using FRNs to evaluate erosion and deposition magnitudes. The aim of this paper is to provide information on application of FRNs in soil erosion, deposition and reservoir sedimentation studies in Africa.

## 2. FALLOUT RADIONUCLIDES (FRNS)

Some of the highly used radionuclides include;  $^{137}\text{Cs}$ ,  $^{210}\text{Pb}$  and  $^7\text{Be}$  (Dercon *et al.*, 2012; Porto *et al.*, 2014a). In applying these nuclides, selection of a reference site is of critical importance. This aids in improving validity of the technique to accurately quantify erosion and deposition. To convert  $^{137}\text{Cs}$ ,  $^{210}\text{Pb}$  and  $^7\text{Be}$  activity to soil erosion, deposition or distribution rate, several models have been developed (Walling *et al.*, 2002; Walling *et al.*, 2014) and they are summarised in **Table 1**. For successful application of the models, a sound understanding of the models and the meaning and derivation of their parameters is highly required to reduce the limitation and uncertainty of information obtained.

### The use of $^{137}\text{Cs}$ in soil erosion and deposition studies

The  $^{137}\text{Cs}$  is an artificial radionuclide which primarily originated from nuclear weapons testing (Ritchie and Mchenry, 1990; Walling, 1998). It was first detected in early 1950s peaking in early 1960s (Stefano *et al.*, 1999; Mabit *et al.*, 2014a) before the nuclear test ban treaty. However, it declined to very low levels in the mid-1970s. The  $^{137}\text{Cs}$  isotope has a half-life of about 30 years (Poręba, 2006). Gellis *et al.* (2006), Walling (1998), Mahawatte and Abeynayake (2010) observed that there is

**Table 1.** Models used in Converting FRNs to Erosion /depositional rates.

Model	Input parameter	FRNs and site used	Advantage	Limitation
Proportional model	Plough depth, bulk density	$^{137}\text{Cs}$ (Cultivated)	1. Requires few input parameters 2. Easy to apply	1. Oversimplification of $^{137}\text{Cs}$ accumulation in the soil 2. Model does not account for dilution of $^{137}\text{Cs}$ concentrations in the soil
Mass Balance Model 1 (MBM1)	Plough depth, bulk density, year of tillage commencement proportional factor, relaxation depth, annual fallout flux*	$^{137}\text{Cs}$ (Cultivated)	1. Accounts for progressive reduction in the $^{137}\text{Cs}$ concentration of the soil	1. Model does not account for removal of freshly deposited $^{137}\text{Cs}$ by erosion during a rainfall event 2. The model assumes that concentration of $^{137}\text{Cs}$ in eroded sediments remains constant through time. Also, the concentration of deposited sediment is assumed to be constant
Mass balance model 2 (MBM2)	Tillage depth, tillage constant, proportional factor, tillage relaxation depth, slope length and slope gradient for each section of the transect, annual fallout flux*	$^{137}\text{Cs}$ (Cultivated) and $^{210}\text{Pb}$ (Cultivated)	1. Accounts for tillage effects 2. Caters for $^{137}\text{Cs}$ dilution in the plough layer 3. Obtained results are more realistic than from the previous models	More data is required than from the previous models
Mass Balance Model 3 (MBM3)	In addition to MB2 parameters Diffusion coefficient, relaxation depth, migration, coefficient, annual fallout flux*	$^{137}\text{Cs}$ (Cultivated) and $^{210}\text{Pb}$ (Cultivated)	1. Results from this model are closer to reality 2. It accounts for tillage redistribution in flow direction	The model is data intensive
Profile shape model	Profile shape factor, Mass depth of soil	$^{137}\text{Cs}$ (Uncultivated) $^7\text{Be}$ (Cultivated)	Simple and easy to use	Does not account for temporal variation of $^{137}\text{Cs}$
Diffusion and migration and model	Diffusion coefficient, Migration rate, /retardation factor	$^{210}\text{Pb}$ (Uncultivated)	1. Accounts of temporal variation of depth profile resulting from the time-dependent fallout input	Requires more time and data to use

\*required for  $^{137}\text{Cs}$  only

a strong adsorption of  $^{137}\text{Cs}$  by clay and organic matter in soils. According to Patrocínio and Andrello (2009) and Dercon *et al.* (2012)  $^{137}\text{Cs}$  is useful in studying the sediment chronology since it is rapidly and strongly absorbed to fine particles of the soil. Also, the distribution of  $^{137}\text{Cs}$  in the soil sediment profile is directly related to the chronology of sediment deposition. Its fallout deposition on land surface from atmospheric nuclear test occurs from rainfall. It is largely absorbed on finer organic and soil particles and it resists leaching through the soil profile (Walling *et al.*, 1995). The profiles of  $^{137}\text{Cs}$  in sediments give the deposition history that can be interpreted to assign calendar dates.

According to Mabit *et al.* (2014d) when assessing the redistribution rates of the soil using  $^{137}\text{Cs}$ , the measured values of  $^{137}\text{Cs}$  in individual sampling points are compared with reference  $^{137}\text{Cs}$  input. The reference inventory represents  $^{137}\text{Cs}$  values measured in undisturbed sites where no erosion nor deposition has taken place (Othman and Ismail, 2012). To reduce the variability of data collected hence ensure that there is uniform local distribution of  $^{137}\text{Cs}$ , reference sites should be selected in the vicinity of the targeted fields that will be subsequently

investigated (Mabit *et al.*, 2013). Selecting appropriate local reference inventory is therefore a major requirement to successful use FRNS in soil erosion and deposition assessment. If the  $^{137}\text{Cs}$  activity is lower than that of reference site, this signifies erosion. On the other hand, if the  $^{137}\text{Cs}$  activity value is higher than that of the reference site, it signifies that soil deposition has taken place (Othman and Ismail, 2012; Porto *et al.*, 2014a). Mostly, in natural uncultivated soils, the distribution of  $^{137}\text{Cs}$  has an exponential decline with soil depth.

According to Walling and Quine (1992), a sharp decline in  $^{137}\text{Cs}$  activity is observed with increasing depth. Minimum vertical translocation of  $^{137}\text{Cs}$  was also reported. In different soils studied, more than 75% of the total inventory occurs in the top 15 cm.

The key assumption that affects application of  $^{137}\text{Cs}$  is that uniform spatial distribution of the nuclide exists (Parsons and Foster, 2011). Local variability of  $^{137}\text{Cs}$  is influenced by individual storm. The application assumes that the patterns from individual storms over extended period leads to uniform local pattern of total fallout (Walling and Quine, 1992). In addition, it is assumed that redistribution of  $^{137}\text{Cs}$  indicates that there is soil move-

ment and that  $^{137}\text{Cs}$  is rapidly adsorbed by soil particles (Mabit *et al.*, 2014a; Walling and Quine, 1992).

Despite these assumptions the use of FRNs provides more information on soil and deposition status in different sites. This method provides an immense potential in data scarce regions. To further improve on the findings the choice of reference sites is critical. Parsons and Foster (2011) reported that use of reference sites in assessing soil erosion and deposition limits the application of FRNs. They argued that samples collected from the same reference sites give high variability of  $^{137}\text{Cs}$ . Another limitation raised by Parsons and Foster (2011), on the use of  $^{137}\text{Cs}$  is its variability in areas affected by the Chernobyl accident, where additional radio-caesium fallout inputs in 1986 can complicate interpretation of  $^{137}\text{Cs}$  measurements. In such areas, use of  $^{210}\text{Pb}$  alone or in combination with  $^{137}\text{Cs}$  in assessing erosion and deposition rates would be highly encouraged.”

Although, possibilities of using  $^{137}\text{Cs}$  in assessing soil erosion and sedimentation in developing countries was discussed in 1995 (IAEA, 1998) very little has been done in most African countries. Since nuclear fallout was low in southern hemisphere regions, it has been noted that  $^{137}\text{Cs}$ , is on decline compared to northern hemisphere (Ritchie and Mchenry, 1990; Walling, 1998; Stefano *et al.*, 1999; Dercon *et al.*, 2012). A great part of sub-Saharan Africa falls on the southern hemisphere. Hence there is difficulty in using  $^{137}\text{Cs}$ , for soil erosion and deposition, since the quantities measured are very low or in some cases below detection limit. As a result, erosion rates can be assessed using a natural radionuclide such as  $^{210}\text{Pb}$ .

### Use of $^{210}\text{Pb}$ in Soil erosion and deposition

This is a natural radionuclide that originates from the decay of  $^{222}\text{Rn}$ , a stable isotope of Radon (Rn) that occurs in soil and the atmosphere. It is a natural radioactive form of lead found in small quantities in most soils as part of uranium decay series and has a half-life of 22.3 years (Lubis, 2006; Porto *et al.*, 2006). According to Guzman *et al.* (2013)  $^{210}\text{Pb}$  is usually derived from radioactive decay of  $^{222}\text{Rn}$  (Radon 222) gas daughter of Radium 226 ( $^{226}\text{Ra}$ ). The  $^{226}\text{Ra}$  occurs naturally in soils and rocks and generates  $^{210}\text{Pb}$  that is in equilibrium with its parent material (Zhang *et al.*, 2014). The soil and sediment samples for measuring  $^{210}\text{Pb}$  is stored for about a month before analysis so as to achieve equilibrium of  $^{226}\text{Ra}$  and its  $^{222}\text{Rn}$  (Shakhashiro and Mabit, 2011; Mabit *et al.*, 2014c). The concentration of  $^{210}\text{Pb}$  is greatest at the surface in uncultivated land because of its continuous replenishment at the soil surface. Also, it has an exponential reduction in activity below the surface (Zhang *et al.*, 2007). The nuclide is useful in establishing the chronology of sediment deposits and estimating the rate of accumulation of such deposits.

The rate of  $^{210}\text{Pb}$  activity declines with depth in a core and it aids in establishing an age depth relationship and

the rate of sediment accumulation. In cases where accumulation rates are high, the  $^{210}\text{Pb}$  activity declines slowly with depth. On the other hand, if accumulation rates are low, the  $^{210}\text{Pb}$  activity reduces rapidly with depth (Walling *et al.*, 2002). The  $^{210}\text{Pb}$  is useful in estimating the soil redistribution for up to 100 years (Zheng *et al.*, 2006; Benmansour *et al.*, 2014; Mabit *et al.*, 2014a) a period longer than that of  $^{137}\text{Cs}$ . Further,  $^{210}\text{Pb}$  is useful in areas where low  $^{137}\text{Cs}$  concentration is likely to pose a problem especially in the southern hemisphere (Porto *et al.*, 2006; Shakhashiro and Mabit, 2011).

A research conducted in different regions under Co-ordinated Research Project (CRP), aimed at investigating the potential of using  $^{210}\text{Pb}$  measurements to estimate soil redistribution rates. It reported that only 31% of the participants provided acceptable results for total  $^{210}\text{Pb}$  analyses (IAEA, 2011). These findings are further supported by Golosov *et al.* (2011) who reported that  $^{210}\text{Pb}$  was characterised by high local variability compared to  $^{137}\text{Cs}$ . Almost similar observation were reported by Mabit *et al.* (2014c) where a major limitation in use of  $^{210}\text{Pb}$  was reported to be the difficulty experienced in accurately analysing the nuclide using gamma spectrometry. There are seasonal variations of  $^{210}\text{Pb}$  in the atmosphere and this affects its use in some parts of the world. Therefore, there is need for more research to test and validate the use of  $^{210}\text{Pb}$  for documenting soil redistribution rates under a range of different environmental and climatic conditions (Benmansour *et al.*, 2014).

Although use of  $^{210}\text{Pb}$  and  $^{137}\text{Cs}$  in erosion assessment over longer periods is possible, their application for short time period is very limited. However, other radionuclides such as  $^7\text{Be}$  nuclide can be used.

### $^7\text{Be}$ use in soil erosion and deposition studies

This is a naturally occurring cosmogenic radionuclide that is produced in the stratosphere and troposphere when oxygen and nitrogen spallate. The nuclide is then deposited as fallout (Guzman *et al.*, 2013). According to Marestoni *et al.* (2009),  $^7\text{Be}$  is a radionuclide that is useful in assessing soil erosion over a short period and in identifying the newly deposited sediment. It is mostly detectable near the surface of the sediment samples and should be counted soon enough after a sample collection since it has a half-life of 53.3 days (Zapata and Agudo, 2000). According to Wilkinson *et al.* (2006), the nuclide is not found below 20 mm of any soil and the radionuclide is useful for surface sediments recovery. The concentration of  $^7\text{Be}$  in the soil is influenced by atmospheric processes, vertical transport in the troposphere and horizontal transport of subtropical and middle latitudes to the tropics and Polar Regions. In addition, it is used in studying mass exchange between troposphere and stratosphere, wet and dry deposition (Marestoni *et al.*, 2009). The main fallout process of  $^7\text{Be}$  is through precipitation which takes 95% of the total. When  $^7\text{Be}$  is used in assessing sedimentation, the stable, eroding and aggrading sites are well

distinguished using both their total inventory and the vertical profile distribution. An exponential decline in activity concentration and areal activity density with depth is usually experienced. Loss of  $^7\text{Be}$  from surface soil indicates that soil erosion has taken place and this removes the upper section of the exponential profile leading to a reduction in overall inventory (Mabit *et al.*, 2014b). The  $^7\text{Be}$  nuclide are useful in assessing the effectiveness of recent soil erosion conservation practices. They are also used to document erosion and sedimentation rates associated with rainfall events. The application of  $^7\text{Be}$  has two key assumptions; the input of  $^7\text{Be}$  associated with the erosional or depositional event is spatially uniform, and, that any pre-existing  $^7\text{Be}$  is uniformly distributed across the area under investigation. Since the method is dependent on individual event, the  $^7\text{Be}$  inventories may be low in arid and semi-arid area of land. Though the nuclide has been used in soil erosion, redistribution and deposition analyses, its application in Africa is very limited.

### 3. APPLICATION OF FALLOUT RADIONUCLIDES IN AFRICA

This section reviews application of FRNs in assessing soil erosion and reservoir sedimentation in Africa. In this, the continent was divided into four regions namely; Eastern Africa, Northern Africa, Western Africa and Southern Africa. During the review, one recent project was found to cover almost all the above-named regions i.e. were well distributed in the African region. The project has been implemented in 10 countries namely; Algeria, Benin, Cote d'ivore, Madagascar, Mali, Morocco, Senegal, Tunisia, Uganda and Zimbabwe. It focused on using Fallout radionuclide's ( $^{137}\text{Cs}$ ,  $^{210}\text{Pb}$  and  $^7\text{Be}$ ) in quantifying African soil degradation, soil erosion and sediment redistribution rates within agricultural landscapes. The

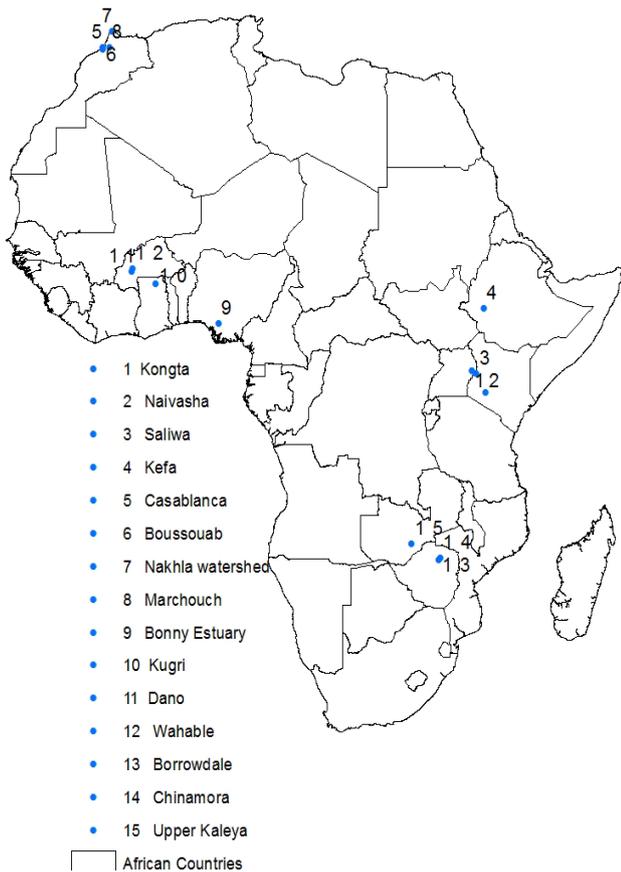
results from this IAEA conducted research project had not yet been published at the period this review paper was compiled. In this paper, the summary of radionuclide activities within the reference site is presented in **Table 2**. In addition, **Fig. 1** shows the distribution of the radionuclide use in Africa. The other findings for different application of FRNs in African region were presented as follows;

#### Eastern Africa

In this region Uganda, Kenya and Ethiopia were found to have some information on use of FRNs in soil and sediment studies. For instance, Ruecker *et al.* (2008) examined the potential of using  $^{137}\text{Cs}$  technique to assess soil redistribution rates in humid tropics of Uganda. The  $^{137}\text{Cs}$  inventory was found to be 392 and 439 Bq/m<sup>2</sup> for Kongta and Magada reference sites, respectively. These results were found to be within the range of globally interpolated  $^{137}\text{Cs}$  inventory of the region. The average  $^{137}\text{Cs}$  inventory within Kongta and Magada was reported to be  $200 \pm 100$  and  $400 \pm 100$  Bq/m<sup>2</sup> respectively. Within the Kongta and Magada sites the mean  $^{137}\text{Cs}$  inventory was reported to be  $202 \pm 87.2$  and  $382 \pm 155$  Bq/m<sup>2</sup> respectively. Further, the net soil redistribution of the two sites (Kongta and Magada) indicated that there was soil loss of  $21 \pm 12.2$  t/ha/yr and  $4.5 \pm 2.3$  t/ha/yr, respectively. Soil deposition within Kongta and Magada was also reported to be  $5.8 \pm 0.9$  and  $6.3 \pm 4.3$  t/ha/yr respectively. It was concluded that  $^{137}\text{Cs}$  method was applicable in estimating soil redistribution within humid tropics of Africa. The study found out that peak hillslope erosion for Kongta and Magada were 44.6 t/ha/yr and 36.3 t/ha/yr while sedimentation rates were 6.4 t/ha/yr and 25 t/ha/yr respectively. The reduced soil deposition and increased soil erosion in Kongta is because the site is on steep hillslope. It was also reported that there had been no documented application of this method in the humid

**Table 2.** Summary of radionuclide activity in reference sites.

Country (Region)	Author	Reference inventory (Bq/m <sup>2</sup> )	FRN used
Uganda (Kongta) (Magada)	Ruecker <i>et al.</i> , 2008	392	$^{137}\text{Cs}$
		439	$^{137}\text{Cs}$
Kenya	Jeff and Shepherd, 2009	<30 (Below minimum detectable limits)	$^{137}\text{Cs}$
Ethiopia (Kefa zone)	Denboba, 2005	$2064 \pm 112$	$^{137}\text{Cs}$
Morocco (Boussouab)	Sadiki <i>et al.</i> , 2007 Benmansour <i>et al.</i> , 2013	1450	$^{137}\text{Cs}$
		3305	$^{210}\text{Pb}$
		1445	$^{137}\text{Cs}$
		101	$^7\text{Be}$
Bukina Faso (Dano and Wahable)	Schmengler, 2011	$729 \pm 14$	$^{137}\text{Cs}$
Ghana	Pennock, 2000	925.1	$^{137}\text{Cs}$
Niger (Southwest)	Chappell <i>et al.</i> , 1998	$2066 \pm 125$	$^{137}\text{Cs}$
Zimbabwe	Quine <i>et al.</i> , 1993	30	$^{137}\text{Cs}$
Zambia (Kaleya)	Walling <i>et al.</i> , 2003 (Collins <i>et al.</i> , 2001)	2602	$^{210}\text{Pb}$
		$202.0 \pm 22.8$	$^{137}\text{Cs}$



**Fig. 1.** Map of Africa showing Regions where FRNs have been used for soil erosion and Reservoir sedimentation analysis.

tropics of Africa. The low application of this method could be due to the region being on the southern hemisphere where  $^{137}\text{Cs}$  inventories are almost one order of magnitude lower than in the northern hemisphere (Walling and He, 1999; Collins *et al.*, 2001; Walling *et al.*, 2002). Ruecker *et al.* (2008) reported difficulty in selecting reference sites near the study area since intensive land use in this climatic zone has been practiced leaving almost all areas being disturbed. They also pointed out that lack of access to institutions with gamma spectrometer detectors in this region further affected widespread applicability of the  $^{137}\text{Cs}$  method.

Further, in Lake Naivasha, Kenya, a sediment core sample was analysed by Stoof-Leichsenring *et al.* (2011) who reported that, total  $^{210}\text{Pb}$  activity reached equilibrium with  $^{226}\text{Ra}$  at a depth of about 33 cm. Also, the unsupported  $^{210}\text{Pb}$  was reported to reduce irregularly with depth but the concentration in the top 15 cm was relatively constant at a mean rate of 164 Bq/kg. Also, at a depth below 30 cm, the concentration was found to be close or near the minimum detection limit. On the other hand, the  $^{137}\text{Cs}$  concentration was found to be highest between depths 20–25 cm. Stoof–Leichsenring *et al.* (2011) also used  $^{210}\text{Pb}$  dating to establish the core chronology that

provided information for a period of 186 yr history of natural and human induced environmental changes.

In addition, Jeff and Shepherd (2009) analysed  $^{137}\text{Cs}$  soil samples in Sauwa river watershed, western Kenya. Four reference sites were selected in the study area and they consisted of sloped grassed pasture, thickly vegetated (Trees and shrubs) and two sites within Kapolete National forest. The sample depth intervals were too broad hence detailed soil profile analysis could not be undertaken. The samples were analysed for  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$  using standard laboratory procedures in IAEA standard reference material (IAEA – 300). The sample counts were taken for 21.7 to 26.1 hours depending on nuclide activity. They reported that most of the samples had low concentration of the nuclide which was well below minimum detection concentration limits of 30 Bq/kg. The  $^{137}\text{Cs}$  was reported to be more concentrated in the upper few centimetres of the soil profile and this concentration decreased with soil depth. Further, little  $^{137}\text{Cs}$  was detected below a soil depth of 20 cm. The mean  $^{137}\text{Cs}$  concentration was reported to be 1.68 and 4.15 Bq/kg in eroded and non-eroded sites. Within the top soils, 82% of samples had detectable level of  $^{137}\text{Cs}$  while 33% had detectable levels within the sub soils. They confirmed the results by field visits that were conducted and 63% of the top soil was found to have been eroded. Almost similar findings were reported for Ethiopia by Denboba (2005).

Denboba (2005) conducted research on the high mountain forests in Kefa zone in southwest of Ethiopia, on the cultivated fields of selected subcatchments. The  $^{137}\text{Cs}$  was used to estimate the rate of soil erosion and deposition. It was found that there was an exponential decline of  $^{137}\text{Cs}$  activity with increase in depth. Further, 69% of  $^{137}\text{Cs}$  inventory was found in the upper 10 cm of soil surface while 19% and 12% inventories were found in 20 cm and 30 cm subsurface layers, respectively. The mean reference inventory was found out to be larger than the assumed range of 200–500 Bq/m<sup>2</sup> (Collins *et al.*, 2001) for the equatorial region. This pointed out the need of more studies within the equatorial region so as to understand the  $^{137}\text{Cs}$  activity other than relying on the world interpolated  $^{137}\text{Cs}$  activity inventories.

### Northern Africa

The  $^{137}\text{Cs}$  was used by Sadiki *et al.* (2007) to quantify erosion rates and identify main factors involved in erosion process in Boussouab catchment, Eastern Rif Morocco. It was found out that at the sampling sites, the  $^{137}\text{Cs}$  activity ranged between 245 and 3670 Bq/m<sup>2</sup>. This indicated that in all sampling sites, erosion had taken place. In addition, soil loss was reported to be highly variable within land use and the variance ranged between 5.1 and 48.8 t/ha/yr. Further, the average soil loss was reported to increase by more than 100% when slope increased from 10° to 25° with loss of 17.7 t/ha/yr and 40.8 t/ha/yr, respectively. The radiometric analysis

proved useful in quantifying erosion rates and in examining the soil movement pattern.

This view is supported by Benmansour *et al.* (2013) in a study conducted in Marchouch South East, Morocco. The findings reported that the net soil erosion rates were 14 t/ha/yr and 12 t/ha/yr when  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$  were respectively used in assessing the soil redistribution in agricultural fields. The study further highlighted that water erosion is the leading process of soil redistribution in cultivated fields. In addition, Benmansour *et al.* (2013) conducted a research on estimating soil redistribution rates and assessment of effectiveness on soil conservation techniques using  $^{210}\text{Pb}$  and  $^{137}\text{Cs}$  isotopes in Marchouch South East from Rabat, Morocco. The net soil erosion using  $^{137}\text{Cs}$ ,  $^{210}\text{Pb}$  and conventional method were reported to be 14.3 t/ha/yr, 12.1 t/ha/yr and 12.3 t/ha/yr, respectively. During the analysis of  $^{137}\text{Cs}$ , Mass Balance Model (MBM<sub>2</sub>) and MBM<sub>3</sub> conversion models were used. The models yielded net soil redistributions of 14.3 t/ha/yr and 15.5 t/ha/yr, respectively. Within the reference site, it was observed that  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$  were mainly distributed in the first 10 cm of the soil depth where the highest concentration was close to the surface and it then diminished with increase in soil depth. They reported that gross erosion and deposition rates in the study area were 12.9 and 0.8 t/ha/yr respectively. The use of combined  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$  FRNs indicated that a potential exists on the use of nuclides in estimating long term erosion rates in agricultural fields in Morocco. The overall data generated on soil redistribution from two nuclides was similar.

Although the concentration of  $^{137}\text{Cs}$  is comparatively low in West African environments, results clearly differentiate zones of sediment loss and gain. The redistribution map of a hillslope in southwestern Burkina Faso shows maximum erosion rates of 16 t ha<sup>-1</sup> yr<sup>-1</sup> at the summit and shoulder position of the hillslope and maximum deposition rates of more than 20 t ha<sup>-1</sup> yr<sup>-1</sup> at the foot slope/valley position (Benmansour *et al.*, 2013)

Benmansour *et al.*, 2014 (in IAEA) assessed long term soil redistribution in a small agricultural field in Morocco using both  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$ . 45 samples were collected along five (5) parallel transects and their activity was measured. The activities were converted into soil redistribution rates using MBM<sub>2</sub>. They found that  $^{210}\text{Pb}$  inventories in the study field ranged between 1700 and 5000 Bq/m<sup>2</sup>. This had a mean relative uncertainty for  $^{210}\text{Pb}$  ranging from 4.6 to 6.6%. The erosion rate was reported to range between 8 and 27 t/ha/yr. Also, in Marchouch Morocco, the  $^7\text{Be}$  was reported by Benmansour *et al.* (2011) in IAEA, (2011) to be successful in quantifying erosion rates during individual or heavy rainfall events. The potential of using  $^7\text{Be}$  in assessing the effectiveness of soil conservation measures was also demonstrated in Morocco in a FAO/IAEA Conducted Research Project (Mabit *et al.*, 2014d). The  $^7\text{Be}$  was found in the first few centimetres of top soil since it has a short half-life and it does not easily move deep down before it de-

grades. Thus, the findings of availability of these nuclides with depth indicate that during radiometric analysis of soil and sediment samples, the top part of soil/sediment should be carefully analysed. In addition, during sample analysis, shorter interval sectioning along the depth should be adopted.

### Western Africa

Schmengler (2011) found out that two cultivated sites at Dano and Wahable in Burkina Faso had  $^{137}\text{Cs}$  areal activity of 200 to 1500 Bq/m<sup>2</sup> and 100 to 2100 Bq/m<sup>2</sup>, respectively. In the two study areas, it was observed that the upper part of the hillslope experienced erosion leading to low  $^{137}\text{Cs}$  inventories while the lower part of the hillslope had higher  $^{137}\text{Cs}$  inventories indicating deposition. Schmengler (2011) reported that the middle part had a balance between erosion process and sediment deposition from the upslope. The net erosion rates for the hill slopes in Dano and Wahable ranged between 2.4–3.3 t/ha/yr and 2.7–5 t/ha/yr respectively. At Bonny estuary in Nigeria, Babatunde *et al.* (2015), studied the different concentration ranges of various radionuclides on sediments. It was found that  $^{137}\text{Cs}$  activity concentration, was lowest in all the study sites. The  $^{137}\text{Cs}$  ranged between 39–67 Bq/kg while  $^{226}\text{Ra}$  ranged between 193–484 Bq/kg. In addition,  $^{238}\text{U}$ ,  $^{232}\text{Th}$  and  $^{40}\text{K}$  were analysed, the results showed a range between 383–812 Bq/kg, 370–819 Bq/kg and 300–9800 Bq/kg for each of the radionuclides, respectively. The elevated  $^{238}\text{U}$ ,  $^{232}\text{Th}$  and  $^{40}\text{K}$  in the region, could be attributed to the presence of cement factories along the banks of the estuary. They reported that cement has high background values of the three nuclides which are associated with limestone, shale, clay, gypsum and other cement raw materials. They reported that the concentration of  $^{238}\text{U}$ ,  $^{232}\text{Th}$  and  $^{40}\text{K}$  in sediment samples was higher than that in raw cement samples of Nigeria. The  $^{137}\text{Cs}$  and  $^{226}\text{Ra}$  radionuclides activity concentrations were not further discussed in the paper and no information exists to link the work to erosion and sedimentation rates. Though the results showed that these radionuclides exist in the sediment core there is need for more investigation to fully understand the results.

Further, measurable quantities of  $^{137}\text{Cs}$  were found in soil samples from profiles in Southwest Niger with a net soil flux of 16 ± 2 t/ha/yr (Chappell *et al.*, 1998). On the other hand, in Kugri farm, Ghana,  $^{137}\text{Cs}$  activity was determined by counting 662 KeV gamma emissions in a period of 2 to 21 hours counting interval. According to Pennock (2000), and (IAEA, 2014) it was reported that though analysis was run on different sized fraction of the concentrations, no measurable  $^{137}\text{Cs}$  was detected. This indicated that, in the region of the study, the  $^{137}\text{Cs}$  would not be applicable in study of soil erosion and deposition and hence other FRNs like  $^{210}\text{Pb}$  should be exploited.

## Southern Africa

An assessment of soil erosion within agricultural land in two contrasting sites (Borrowdale and Chinamora) near Harare, Zimbabwe was conducted by Quine *et al.* (1993) using  $^{137}\text{Cs}$ . The  $^{137}\text{Cs}$  technique was reported to be applicable in studying soil erosion within the region. Borrowdale area was characterised by commercial farming and tropical red soils while Chinamora area had communal farming and granite sandy soils. Within the undisturbed regions of both study sites, the reference inventory was found to be  $30 \text{ Bq/m}^2$  while the profile distribution on cultivated soils showed a mixed  $^{137}\text{Cs}$  throughout the plough layer. In addition, applicability of  $^{210}\text{Pb}$  technique in assessing soil loss and deposition under different land uses was investigated by Collins *et al.* (2001) and Walling *et al.* (2003) in upper Kaleya River basin, Zambia.

Collins *et al.* (2001) used  $^{137}\text{Cs}$  for different land uses. The considered land uses included; commercial cultivation, communal cultivation and bush grazing. The sample count time of 50,000s was used due to low activity of nuclide. During the study, general guidelines and criteria by Walling and Quine (1993) for identifying representative reference sites were followed. 36 core samples were collected in sites with minimum slopes, fully vegetated and where no deposition and erosion has taken place. Sampling was stratified along the transect to cover areas below where communal cultivation, commercial cultivation and bush grazing land use were practiced. The theoretical MBM for  $^{137}\text{Cs}$  was used. The gross erosion, net soil loss and deposition within communal cultivated site reported to be 7.0, 2.5 and 4.5 t/ha/yr while that of commercial cultivation was 7.0, 4.3 and 2.7 t/ha/yr respectively. In bush grazing site, gross erosion was lowest at 3.4 t/ha/yr and net soil loss and deposition was reported to be 2.9 and 0.5 t/ha/yr. These results were based on transects sampled within areas characterised by a given land use and not on the entire topographic unit (slope to river channel)

Walling *et al.* (2003) compared  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$  in estimating the net soil redistribution under different land uses in upper Kaleya River basin, Zambia. The results from commercial cultivation gave soil loss amounts of 4.3 t/ha/yr and 4.5 t/ha/yr while those of communal cultivation were 2.5 t/ha/yr and 1.2 t/ha/yr, respectively. Further, bush grazing had a net soil redistribution of 2.9 t/ha/yr and 3.7 t/ha/yr respectively. In the study using  $^{210}\text{Pb}$ , soil deposition was found to be 3.7, 5.2 and 4.5 t/ha/yr in commercially, communal cultivated and bush grazing land uses respectively. The findings by Walling *et al.* (2003) are comparable with those of Collins *et al.* (2001) who used  $^{137}\text{Cs}$  within the same catchment. The  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$  has reported similar trend in gross erosion, net soil loss and deposition assessment. Also, it was reported that unsupported  $^{210}\text{Pb}$  can be used as an alternative to  $^{137}\text{Cs}$  in equatorial region of Africa where  $^{137}\text{Cs}$  inventories are very low.

## 4. CONCLUSION

A challenge still exists in selecting suitable reference sites for radionuclides. There is need for more actual measurements of erosion using different methods to improve on the quality of results. In addition, the reference sites should be selected as close as possible to the areas erosion and deposition is being assessed. Although use of FRNs have been reported to be useful in assessing soil redistribution and reservoir sedimentation in different regions of world, there is a long way to go in applying FRNs in Africa. Hence more research is required to exploit the use of FRNs for soil erosion and reservoir sedimentation analysis. Some of the FRNs that have been used include;  $^{137}\text{Cs}$ ,  $^{210}\text{Pb}$  and  $^7\text{Be}$ . Most of these radionuclide activities in Africa have been found detectable up to a maximum depth of 30 cm. Morocco in Northern Africa is one of the country in which a lot has been done from 2000s' in the use of FRNs in soil erosion and soil redistribution assessment. In the southern Africa, high  $^{137}\text{Cs}$  activity was reported back in 1993. This radionuclide activity could have diminished with time hence the need for a more current study.

In the summarised review on African continent, most studies that have been conducted, are on use of  $^{137}\text{Cs}$  fallout radionuclide. In some cases, comparative study has been undertaken using  $^{137}\text{Cs}$  and  $^{210}\text{Pb}$ . The results from the reference inventory measured, indicated that  $^{210}\text{Pb}$  activity is higher than the  $^{137}\text{Cs}$  activity. Also in some regions, the  $^{137}\text{Cs}$  activity was reported to be below the minimum detection limit implying that the technique cannot be used in assessing soil erosion and deposition and that more studies/research are required for the region. Further, the highest  $^{137}\text{Cs}$  activity in the reference site has been reported in Niger at a value of  $2066 \pm 125 \text{ Bq/m}^2$  a region that is in the northern hemisphere which received more of  $^{137}\text{Cs}$  activity than southern hemisphere.

Though the application of  $^{137}\text{Cs}$  in the southern hemisphere is minimal, due to its low  $^{137}\text{Cs}$  activity, the use of other fallout radionuclides such as  $^{210}\text{Pb}$  and  $^7\text{Be}$  need to be explored. As a result, more studies are required in Africa on application of  $^{210}\text{Pb}$  in areas that would be representative of the whole continent to better understand soil erosion processes and reservoir sedimentation. Also,  $^7\text{Be}$  nuclide is least applied within Africa. Thus, to understand short term effects on soil loss/redistribution,  $^7\text{Be}$  should be analysed. To improve on assessing soil loss, redistribution and deposition analysis in Africa, FRNs should be highly exploited. From the reviewed literature on Africa use of FRNs, only in Lake Naivasha Kenya, have they been applied to study sedimentation and sediment chronology. This indicates that there is need to link use of FRNs in the erosion and deposition studies within the catchment to sedimentation of reservoirs. This can be achieved by analysing more sediment cores from reservoirs to generate sedimentation rates and chronology.

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