



TESTING THE STANDARDIZED GROWTH CURVE (SGC) TO OSL DATING COASTAL SEDIMENTS FROM THE SOUTH BOHAI SEA, CHINA

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Abstract: It has been suggested that the standardized growth curve (SGC) method can be used to accurately determinate equivalent dose (D_e) and reduce measurement time. However, different opinions regarding the applicability of the SGC method exist. In this paper, we evaluated quartz OSL SGCs of marine and coastal sediments of different grain sizes and different cores in the south Bohai Sea in China, and tested their applicability to the determination of D_e values. Our results suggested as follows: (1) The SGC method is applicable to both multiple- and single- aliquot regenerative-dose (MAR and SAR) protocols of OSL dating and efficiently provides reliable estimates of D_e . (2) Fine-sand quartz of different palaeodoses showed highly similar dose-response curves and an SGC was developed, but old samples using the SGC method have large uncertainties. (3) For coarse-silt quartz, two different types of dose-response curves were recorded: low-dose (≤ 60 Gy) and high-dose (≥ 100 Gy). The growth curves of low-dose quartz were similar to each other, facilitating the use of SGC in D_e estimations, but errors tended to be larger than those obtained in the SAR method. For high-dose (100-300Gy) quartz, the SGC was also found to be reliable, but there was large uncertainty in D_e (>300 Gy) estimation. We suggest that SGC could be employed for the dating of marine and coastal sediments dating using either MAR or SAR OSL protocol and either fine-silt, coarse-silt or fine-sand quartz.

Keywords: Standardized growth curve (SGC), OSL dating, marine sediments, Bohai Sea in China.

1. INTRODUCTION

Coastal areas are some of the world's most populated and dynamic landscapes on Earth. Understanding their formation and assessing the past, present, and future impact of change in their geological evolution of our coasts requires reliable chronological tools. The potential and desirability of luminescence dating methods in the determination of the magnitude and timing of events recorded by coastal deposits have long been recognized (see a reviews by Jacobs, 2008), but the lack of precision and the uncertainties in application have prevented them from making breakthrough contributions to various issues (e.g. Murray and Olley, 2002). Recent progress in both methodology and technology now allows demonstration and explicit testing of the reliability of the calculations, making luminescence dating a very attractive and increasingly used approach.

Standardized growth curve (SGC) in optically stimulated luminescence (OSL) dating was first proposed by Roberts and Duller (2004), as a means of accurately estimating equivalent dose (D_e) through a simplified procedure and thus speeding up the measurement process. When SGC is used in a single aliquot regenerative-dose (SAR) protocol (Murray and Wintle, 2000), the SGC of a sample can be determined based only on measurements of natural luminescence signal intensity (L_n) and its response to artificial irradiation dose (T_n) which speeds up the OSL dating process considerably (Roberts and Duller, 2004). The applicability of SGC has since been tested in sediments from agricultural areas (e.g. Burbidge *et al.*, 2004), loess and desert (e.g. Lai, 2006; Lai *et al.*, 2007; Stevens *et al.*, 2007; Yang *et al.*, 2011; Kang *et al.*, 2012), floodplains (e.g. Shen and Mauz, 2011); coastal dunes (e.g. Telfer *et al.*, 2008), and lacustrine deposits (e.g. Long *et al.*, 2010).

These studies have provided various perspectives on the applicability SGC to OSL dating. Lai *et al.* (2007) proposed that a universal SGC for different continental loesses exists only when D_e is lower than 200 Gy. Although good consistency was observed, the contamination of heavy minerals in loess samples (Lai *et al.*, 2007) or feldspar in lacustrine samples (Long *et al.*, 2010) should be considered carefully during an SGC application. Telfer *et al.* (2008) reported that the SGC is a regionally based method and only performs well when determining low (<50 Gy) D_e values. It has also been reported that the SGC approach is valid only when an SGC is constructed and applied to samples from the same section in Chinese loess (Stevens *et al.*, 2007) or pottery pieces (Hong and Choi, 2008). As a conservative position, Lai (2006) reported various performances for young (<270 ka) and old samples of Chinese loess, and Yang *et al.* (2011) constructed two SGCs for desert samples with low (<80 Gy) and high D_e values collected from northern China. From these works, it can be concluded that SGC is applicable to the determination of D_e in rapid OSL measurements, but

its performance varies by sample type and study area. So far, little attention has been paid to marine and coastal sediments.

Recently, luminescence dating has been widely applied to palaeoenvironmental studies around the Bohai Sea, China. These include assessments of tectonic activity (e.g. Xu *et al.*, 2005), alternations between transgression and regression (e.g. Chen *et al.*, 2008; Yi *et al.*, 2012a), environmental reconstructions (e.g. Han, 2001; Yi *et al.*, 2012b) and human activities (e.g. Zhang *et al.*, 2010). However, the parameters of the existing SGCs varied with the types of sediment and continents, and large offset was observed when previous results were used in estimating D_e s of sediments collected from the Bohai Sea, suggesting different OSL characteristics of quartz of this type sediment or of this area. Thus, it is necessary to assess the performance of the SGC method in determining the luminescence characteristics of coastal sediments. In this paper, we focus on the applicability of the SGC method to OSL dating using marine and coastal sediments believed to date to the late Pleistocene. The samples used in this study were collected from three boreholes drilled in the south Bohai Sea, China. The test of applicability was conducted from internal comparison to external comparison across sections. Three general SGCs were obtained for quartzes of various grain sizes.

2. MATERIALS AND MEASUREMENT TECHNIQUES

Regional settings

The south Bohai Sea (Laizhou Bay, Fig. 1) is located between branches of the Yi-Shu Rift (Gao *et al.*, 1980; Zhang *et al.*, 2003). The period from the Neogene to the present has been marked by tectonic quiescence and stable sedimentation rate (Wu *et al.*, 2006; Yu *et al.*, 2008). Sedimentary alternations were mainly between delta, estuarine and tidal plain systems (Xue and Ding, 2008). During regressions, the exposed area of the south Bohai Sea would have been replaced in part by diluvial fan (Chen *et al.*, 1991), loess/sandy dune (Chen *et al.*, 1991; Zhao, 1991, 1995; Yu *et al.*, 1999) or alluvial fan (Meng *et al.*, 1999). The drainage system contains Xiaoqinghe River, Mihe River and Weihe River, which all originate from the Luzhong Mountain Range with several hundreds of kilometres.

Samples for OSL dating

Samples from shallow sea to deltaic environments for OSL dating were collected from three boreholes, i.e. Lz908, BH1 and BH2, drilled in the south Bohai Sea by First Institute of Oceanography, State Oceanic Administration, China. Borehole Lz908 (37°09'N, 118°58'E, elevation 6 m) was drilled in the summer of 2007. The drilling location was covered by seawater until the middle of last century. The core's length is 101.3 m with a diam-

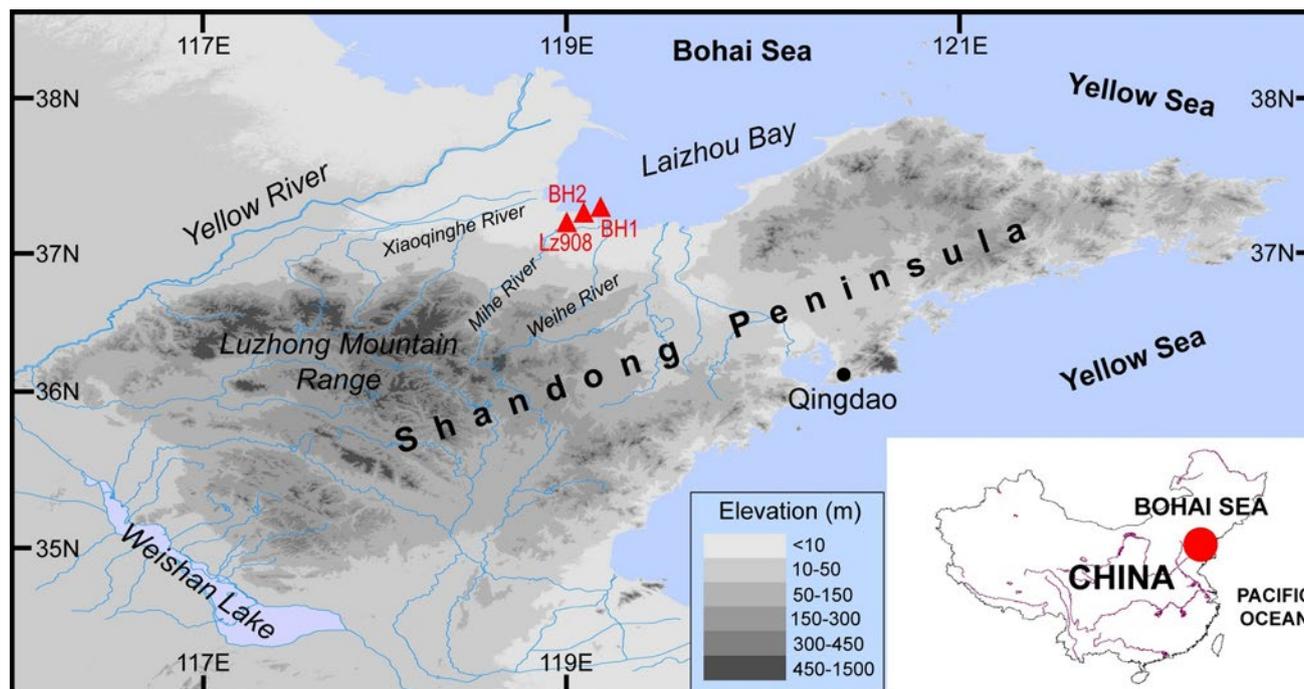


Fig. 1. Location of the south Bohai Sea in China, Geographical settings and position of Lz908, BH1 and BH2 cores (▲).

eter of 110 mm and the recovery rate is 75%. Borehole BH1 (37°17'N, 119°06'E, water depth -4 m) was drilled in the winter of 2008. The core's length was 198.8 m and the recovery rate was 90%. Borehole BH2 (37°10'N, 119°04'E, elevation 3 m) was drilled in the spring of 2009. The drilling position was covered by seawater until the middle of last century. The core's length was 228.2 m and the recovery rate was 85%. All of these samples were well dated (Yi *et al.*, 2012a, 2012b) and displayed similar luminescence characters allowing testing the SGC method to OSL dating coastal sediments.

Sample pre-treatment and measurement techniques

Lz908 core

Nine OSL samples were taken by hammering metal cylinders of 2.5 cm (diameter) × 6 cm (length) horizontally into the core after being cut half. The samples were extracted under subdued red light and pretreated with 30% HCl and 30% H₂O₂ to remove the carbonates and organic material, respectively. Then the fine silt (4-11 μm) was obtained using sedimentation procedures based on Stokes' Law. These polymineral fine grains were immersed in H₂SiF₆ (30%), for a week in an ultrasonic bath to obtain the fine-grained quartz component. The purity of prepared quartz was tested by the routine IR stimulation (Lai and Brückner, 2008), and then were mounted on the centre part with a diameter of ~0.5 cm. In our discs the grains are sparsely distributed and about 30000-50000 grains could be mounted according to Duller (2008) on stainless steel discs of a diameter of

1 cm. For D_e determinations, we chose the pure quartz of fine silt (4-11 μm) and followed the sensitivity-corrected multiple aliquot regenerative-dose protocol (MAR) developed by Lu *et al.* (2007) to determine the D_es. All measurements were performed using a Daybreak 2200 automated OSL reader in Qingdao Institute of Marine Geology, Chinese Geological Survey. Stimulation used blue LED (λ=470±5 nm) for 50 s at 125°C. The OSL signal was detected by an EMI 9235QA photomultiplier tube and two 3 mm U-340 glass filters. The MAR conditions were: preheat at 260°C for 10 s, cut-heat at 220°C for 10 s; test doses (TD) of 1, 6, 10, 15, and 25 Gy were used for different samples, respectively; for each sample, one regeneration dose, ranging from 3 to 320 Gy according to the magnitude of the natural dose of that sample, were used for the construction of the growth curves.

BH1 and BH2 cores

OSL samples were taken by hammering metal cylinders of 5 cm (diameter) × 15 cm (length) vertically into the cores before being cut. Sample preparation was done under subdued red light in OSL laboratory. All samples were treated with 10% HCl and 30% H₂O₂ to remove carbonates and organics, respectively. We choose twenty-four samples of coarse-silt grain (38-63 μm) and nine samples of fine-sand grain (60-90 μm), when available, for OSL measurements. Grain-size fractions of 38-63 μm and 63-90 μm were sieved. Quartz fraction of silty grains (38-63 μm) was then treated with silica-saturated fluoro-silicic acid (H₂SiF₆) for about two weeks (Berger *et al.*, 1980; Roberts *et al.*, 2003) to remove feldspars. For fine-

sand grains (63-90 μm), we used hydrofluoric acid (HF) for 45-60 minutes to remove feldspars. The purity of prepared quartz was tested by the routine IR stimulation (Lai and Brückner, 2008). The quartz grains were then mounted on the centre part with a diameter of ~ 0.5 cm. In our discs the grains are sparsely distributed and about 3000-5000 and 300-500 grains for silty and fine-sandy quartz, respectively, could be mounted according to Duller (2008) on stainless steel discs of a diameter of 1 cm.

OSL measurements were carried out using a Risø DA-20 automated TL/OSL system at the Luminescence Dating Laboratory of Qinghai Institute of Salt Lakes, CAS. Stimulation used blue LED ($\lambda=470\pm 20$ nm) for 40 s at 130°C. 90% diode power was used. OSL signal was detected using a 9235QA photomultiplier tube through a 7.5 mm thick Hoya U-340 detection filter.

D_e s were determined using the SAR protocol method (Murray and Wintle, 2000). To select an appropriate preheat temperature for D_e determination, preheat plateau test was conducted for two samples, BH1-3B and BH2-16. Preheat temperatures from 220 to 300°C (for 10 s) with an interval of 20°C were tested, using a cut-heat of 220°C for 10 s. A good preheat plateau was observed from 220 to 280°C for the two samples. So, a preheat temperature of 260°C was selected for D_e determination. These two samples were also taken for dose recovery test. The ratios of the given doses to the measured doses were 0.95 ± 0.03 and 0.99 ± 0.02 , respectively. This result suggested that the SAR procedure is able to recover a laboratory dose. Thus, preheat was chosen to be at 260°C for 10 seconds and cut-heat at 220°C for 10 s. OSL signals from the first 0.64 s stimulation were integrated for growth curve construction after background subtraction. For each sample, six or four regeneration doses, ranging from 0 to 600 Gy, were used for the construction of the growth curves, whose test dose was set as 11.1 Gy for all samples. All aliquots satisfied the prerequisite that the recycling ratio is to be within the range of 0.9-1.1.

3. RESULTS AND DISCUSSION

According to Roberts and Duller (2004), the SGC process could be described as follows: (1) A luminescence signal (L_x) is normalized using a test-dose response (T_x); (2) The normalized luminescence signal (L_x/T_x) is multiplied by the test dose (T_D); (3) An SGC from the regenerative dose (R_D) is fit to the standardized luminescence signal, i.e. $((L_x/T_x)*T_D)$; (4) The D_e of the sample is determined based only on measurements of the normalized natural luminescence signal $((L_n/T_n)*T_D)$. The shape of fitting formula can be expressed as an exponential form (e.g. Burbidge *et al.*, 2004; Roberts and Duller, 2004; Stevens *et al.*, 2007; Hong and Choi, 2008; Telfer *et al.*, 2008) or as an exponential form plus a linear component (e.g. Lai, 2006; Lai *et al.*, 2007; Long *et al.*, 2010;

Yang *et al.*, 2011). An exponential plus linear form is described as follows:

$$Y = k \times (1 - e^{-\frac{X}{D_0}}) + c \times X + d \quad (3.1)$$

Here

$$Y = \frac{L_x}{T_x} \times T_D \quad (3.2)$$

is standardized OSL, X is the regenerative dose, k is the standardized OSL intensity at saturation level for exponential growth part, D_0 is the characteristic dose at which the slope of growth curve (exponential part) is equal to $1/e$ of the initial value, and c and d are constants defining the linear component. We choose Eq. 3.1 as the target function for testing of the SGC method.

The sedimentary processes that the grains experienced are important factors that affect grains' luminescence characteristics. For instance, the characteristics of loess, such as its medium-to-long transport distances, suggest that any previous luminescence signal should be removed prior to deposition, making it an ideal material for the testing of luminescence techniques (Roberts, 2008). However, the complex hydrodynamics of marine and coastal sediments can result in large errors in estimations of luminescence ages (Jacobs, 2008). Grains that settled in different positions also experience different sedimentary processes. As stated by Chen *et al.* (in press), in the southern Bohai Sea, silty grains are suspended and move seaward whereas fine-sand grains are trapped *in situ*. For same grains, sedimentary alternations were mainly between deltaic, estuarine and tidal plain systems (Xue and Ding, 2008). A few tens of sandy quartz grains can significantly vary dose responses (Duller, 2008). We therefore tested the suitability of SGC in dating coastal sediments starting with comparisons of the D_e s of aliquots, of samples, and of cores.

SGC of fine-silt (4-11 μm) quartz extracted from the Lz908 core

Standardized growth curve construction for fine-silt quartz

Because of the dependence of the standardized dose response upon the size of test dose (Shen and Mauz, 2011), a standardized OSL obtained for fine-silt quartz was first converted to an equivalent standardized OSL, according to the calibration developed in Shen and Mauz (2011). Combining the data from all nine samples, an SGC can be produced by following the steps proposed by Roberts and Duller (2004) (Fig. 2):

$$Y = 37.186 \times (1 - e^{-\frac{X}{56}}) + 0.134 \times X + 0.754 \quad (3.3)$$

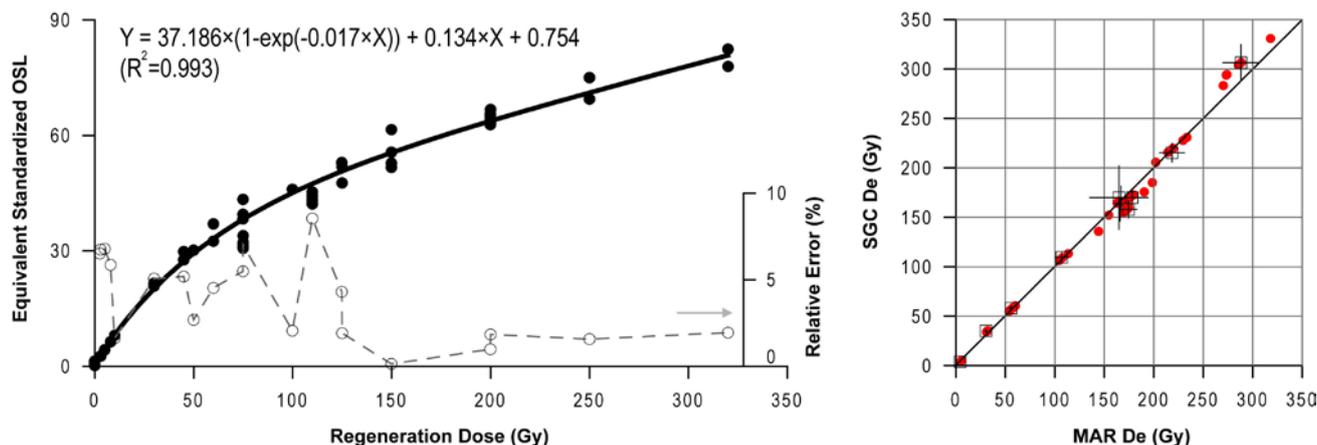


Fig. 2. SGC constructed using the fine-silt quartz of nine samples from Lz908 core (left panel). Also shown is the fitting equation, together with R^2 value. The recuperation is negligible for these samples, as the fitting line almost goes through the original point. In all cases, the error bar is smaller than the symbol (right-hand y-axis). Comparison of $D_{e,s}$ between MAR and SGC methods is displayed in right panel. Solid cycles in the right panel indicate individual $D_{e,s}$.

The explained variance $R^2=0.990$, the freedom-adjusted explained variance $R^2_{adj}=0.993$, the standard deviation of the regression $s.e.=2.246$, and all coefficients are significant at $p<0.01$.

Comparison between $SGC_{fine-silt}$ and MAR

Because the individual growth curve is constructed from all aliquots and because the fitting domain of a given SGC is exactly same as the individual growth curve produced using an MAR protocol, we cannot directly compare the $D_{e,s}$ of $SGC_{fine-silt}$ and individual growth curves. For this reason, we developed a substituted process to assess the applicability of SGC: (1) An SGC_i is determined using eight out of the nine samples; (2) This SGC_i is used to estimate the $D_{e,s}$ of the remaining sample; (3) The $D_{e,s}$ produced using the MAR and SGC methods are compared; (4) These steps are repeated nine times, leaving a different sample out of step (1) each time.

A graphic comparison between $D_{e,s}$ produced using the SGC and MAR methods is displayed in **Fig. 2**. The averaged $D_{e,s}$ of five samples were found to match each other well. The SGC method was found to produce smaller D_e values than the MAR method for three samples and larger D_e values for two samples. The standard deviation values of SGC method were all significantly smaller than those of the MAR method.

A statistical comparison between $D_{e,s}$ from SGC and MAR methods was conducted using one-way analysis of variance (ANOVA) which is used to test the hypothesis that several means are equal (Kleinbaum *et al.*, 1997). Herein, the dependent variables are $D_{e,s}$, and the single-factor (independent) variable is the method, SGC or MAR. Results are displayed in **Table S1**. They indicate that the $D_{e,s}$ of most of the samples were not significantly different between the two methods. It can be concluded that, for 4-11 μm quartz samples from Lz908 core, $D_{e,s}$

calculated using the SGC method are statistically the same as those calculated using the MAR method, and SGC is suitable for use in the OSL dating of fine-silt quartz OSL dating.

Fine-sand (63-90 μm) samples processed using the SAR protocol

Comparison of growth curves of individual aliquots of the same sample

Figs. 3A-D shows individual growth curves of six aliquots for samples BH1-6B, BH1-14B, BH2-5 and BH2-15, respectively. The growth of the normalized OSL signal with regeneration dose is very similar across all six aliquots within a sample. This pattern was observed for all samples dated using the SAR protocol in this study.

Comparison of growth curves of individual samples from the same core

The average growth curves of all nine fine-sand samples from BH1 and BH2 cores are shown in **Figs. 3E-F**, respectively. The normalized OSL signal levels for a given regeneration dose are identical, with small standard errors. This implies that the samples from same core shared the same dose-response curve.

Comparison of growth curves from two cores

In order to examine the statistical fidelity of a SGC in different dataset, two average growth curves from BH1 and BH2 cores were estimated. Results are shown in **Figs. 3E and F**, respectively. Each curve was then used to estimate the D_e of the other curve. The parameters of these two growth curves were very similar, and the D_e offsets were below 20%, suggesting the independence of the SGC method on different cores.

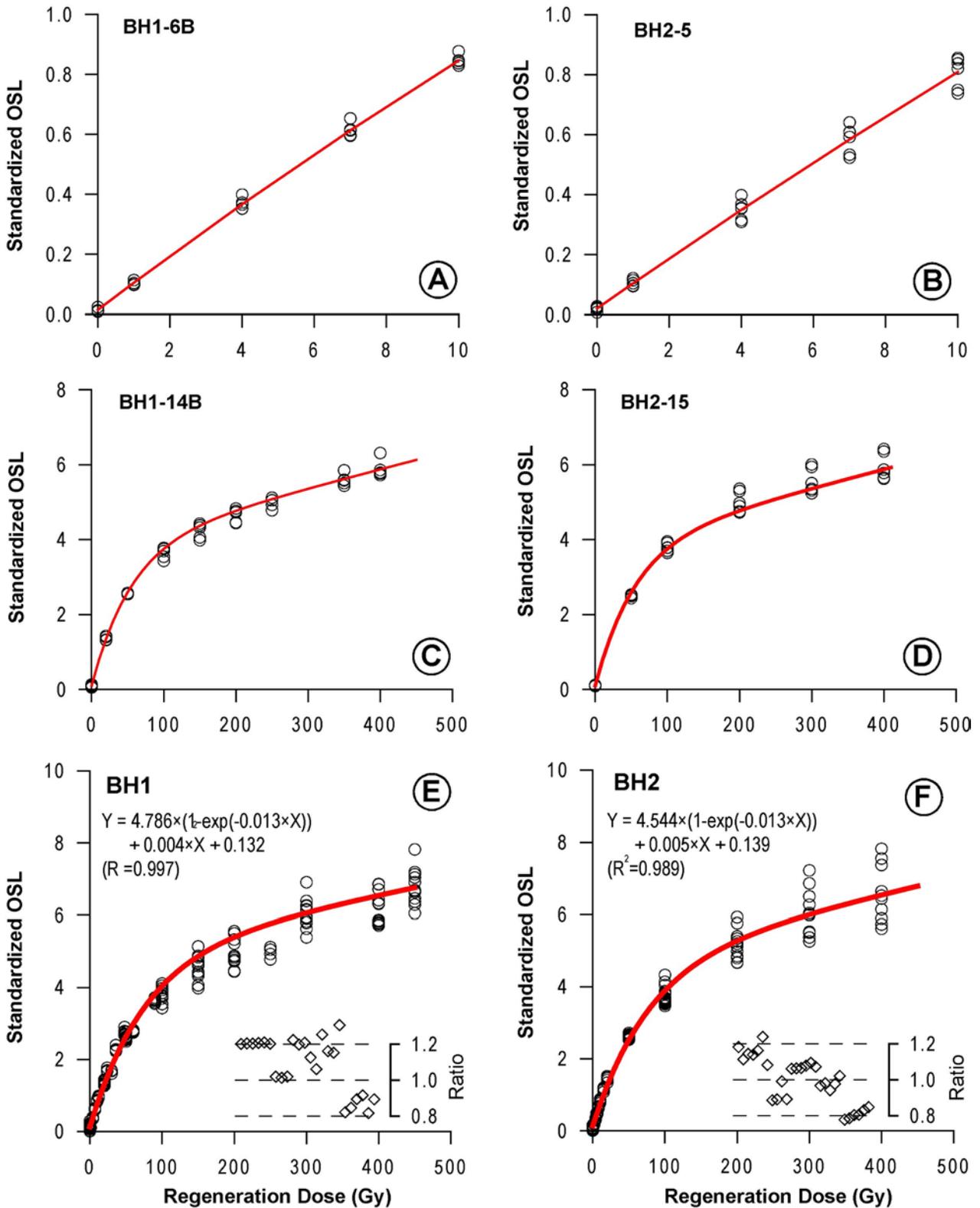


Fig. 3. (A-D) Growth curves of six aliquots for sample BH1-6B, BH2-5, BH1-14B and BH2-15, respectively. The bold lines represent the average curves of each sample. (E-F) Average growth curves (bold lines) of all BH1 and BH2 samples, respectively, estimated from averaging standardized OSL data within samples. Inserts in E and F show ratios between D_0 s based SAR and SGC methods: (1) Estimate a SGC based on BH1 dataset, and then estimate D_0 s of BH2 dataset. The results are displayed in E. (2) Estimate a SGC based on BH2 dataset, and then estimate D_0 s of BH1 dataset. The results are displayed in F. Two equations are displayed in figures, indicating similarity of growth curves between cores. D_0 offsets are smaller than 20%, suggesting independence of SGC method on different cores.

Standardized growth curve construction for fine-sand quartz

The comparisons of growth curves in samples, in cores and between cores described above demonstrated that it is possible to construct a regional SGC for fine-sand quartz. An SGC was estimated by combining the data from all nine samples as below (Fig. 4):

$$Y = 4.286 \times (1 - e^{-\frac{X}{63}}) + 0.005 \times X + 0.106 \quad (3.4)$$

The explained variance $R^2=0.998$, the freedom-adjusted explained variance $R^2_{adj}=0.997$, the standard deviation of the regression $s.e.=0.110$, and all coefficients are significant at $p<0.01$.

As shown, all the regeneration dose points conform closely to the fitting line. The constant, here 0.106, implies that the recuperation for these samples is negligible. This shows that the construction of the $SGC_{fine-sand}$ was acceptable.

Comparison of D_e s estimated using $SGC_{fine-sand}$ and SAR

For evaluation, we used $SGC_{fine-sand}$ to determine the D_e s of the nine samples of BH1 and BH2 cores. A graphic comparison between D_e s from $SGC_{fine-sand}$ and SAR is shown in Fig. 4. The averaged D_e s of the four samples matched each other closely. SGC produced smaller D_e values than SAR for three samples, and larger D_e values for two samples. The standard deviations of the two methods were similar.

A statistical comparison between D_e s from $SGC_{fine-sand}$ and SAR methods was conducted using the One-Way ANOVA. The results (Table S1) showed a significant difference in the two samples, specifically BH1-17B and BH2-13. When a few dozen quartz grains were measured, a significant dose response variation was observed (e.g.

Duller, 2008). This caused the differences in the D_e s of these two samples. Because many studies showed that the uncertainties of dating coastal and marine sediments were significantly larger than those of dating other deposits (e.g. Jacobs, 2008), the observed difference could also be related to the sedimentary process. It can be concluded that an SGC can be used in determination of the D_e values of fine-sand quartz, but a careful assessment is necessary in the dating of older sediments.

Coarse-silt (38-63 μ m) samples processed using the SAR protocol

Comparison of growth curves of individual aliquots of the same sample

Figs. 5A-D show the growth curves of BH1-3B, BH1-11A, BH2-4B and BH2-16, respectively. The growth of the normalized OSL signal with regeneration dose was found to be similar for all aliquots within a given sample. This pattern was seen for all samples which have been dated using the SAR protocol in this study.

Comparison of growth curves of individual samples from same core and between cores

Figs. 5E-F show average growth curves of all coarse-silt samples of the BH1 and BH2 cores, respectively. In SGC applications in fine-silt and fine-sand quartz grains, the offset between low- and high-dose samples was small (Figs. 2-4), but the offset was highly visible in coarse silty quartz grains (see insets of Figs. 5E-F).

To determine the dose boundary for the low- and high-dose samples, we created a scatter plot of regeneration dose changes against the ratio between D_e s based on the SAR and SGC methods (Fig. 6A). The figure shows that the ratio varied considerably when the dose was below 100 Gy and clustered around 1 when the dose

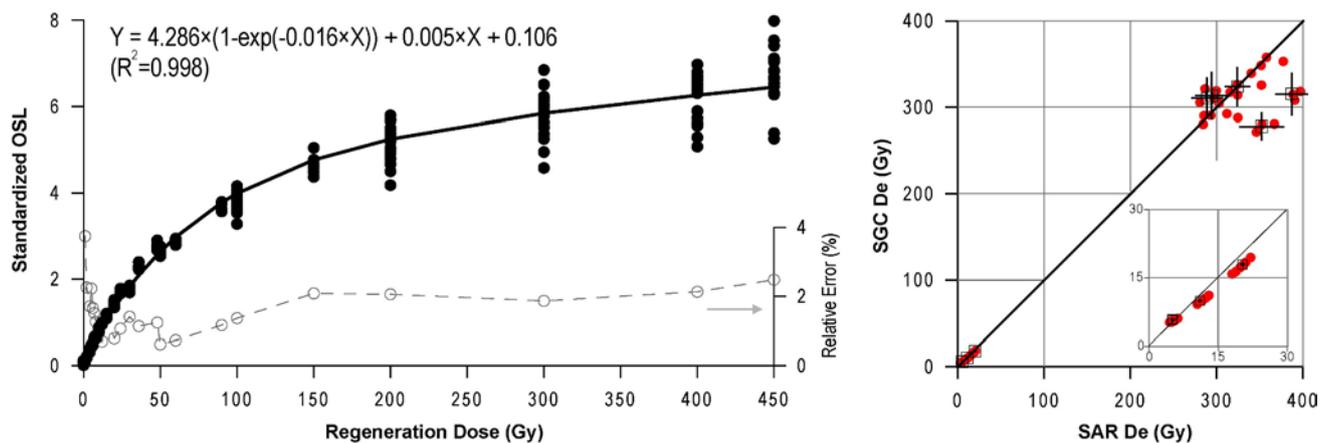


Fig. 4. The SGC constructed using the fine-sand quartz of nine samples from BH1&BH2 cores, estimated from averaging standardized OSL data within samples (left panel). Also shown is the fitting equation of the SGC, together with R^2 value. The recuperation is negligible for these samples, as the fitting line almost goes through the original point. In all cases, the error bar is smaller than the symbol (right-hand y-axis). Comparison of D_e s between SAR and SGC methods (right panel). Solid cycles in the right panel indicate individual D_e s.

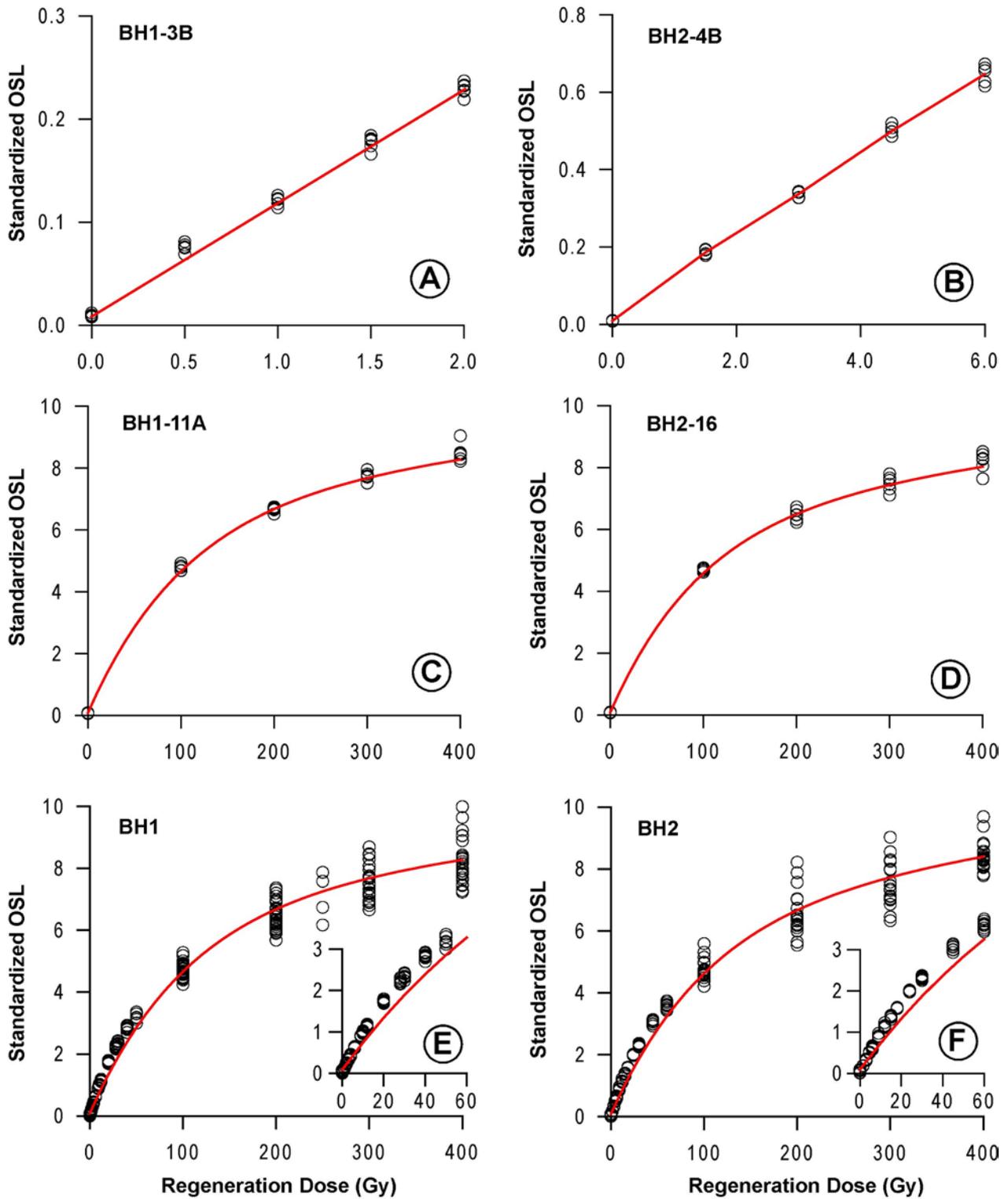


Fig. 5. (A-D) Growth curves of six aliquots for sample BH1-3B, BH2-4B, BH1-11A and BH2-16, respectively. The bold lines represent the average curves of each sample. (E-F) Average growth curves (bold lines) of all BH1 and BH2 samples, respectively, estimated averaging standardized OSL data within samples. The difference for relative low doses (≤ 60 Gy) is significantly under-estimated.

was higher than 100 Gy, indicating a clear boundary at 100 Gy. Considering that there was no sample with a dose between 60 and 100 Gy, we divided these twenty-four samples into two groups: low-dose (≤ 60 Gy) and high-dose (≥ 100 Gy).

The results of several previous studies of SGC method showed different levels of performance for young and old samples and the boundary was found to vary with types of sediment. For instance, Telfer *et al.* (2008) found the boundary to be 50 Gy, Yang *et al.* (2011) found it to be 80 Gy, and it was found to be 270 Gy by Lai (2006).

However, regardless the type of sediment, the offset between low- and high-dose samples indicated differences in the luminescence characteristics of young and old samples. The cause of this difference can be sedimentary process or D_e 's variation. Yi *et al.* (2012c) detailed four grain-size components of sediment in this area, classifying them as those whose modes were finer than 4, 4.6-12.5, 23.4-63.3 and 67.1-132.6 μm , respectively. They stated that the sedimentary process of the large two components did not change significantly. The offset between low- and high-dose coarse silty quartz could not have

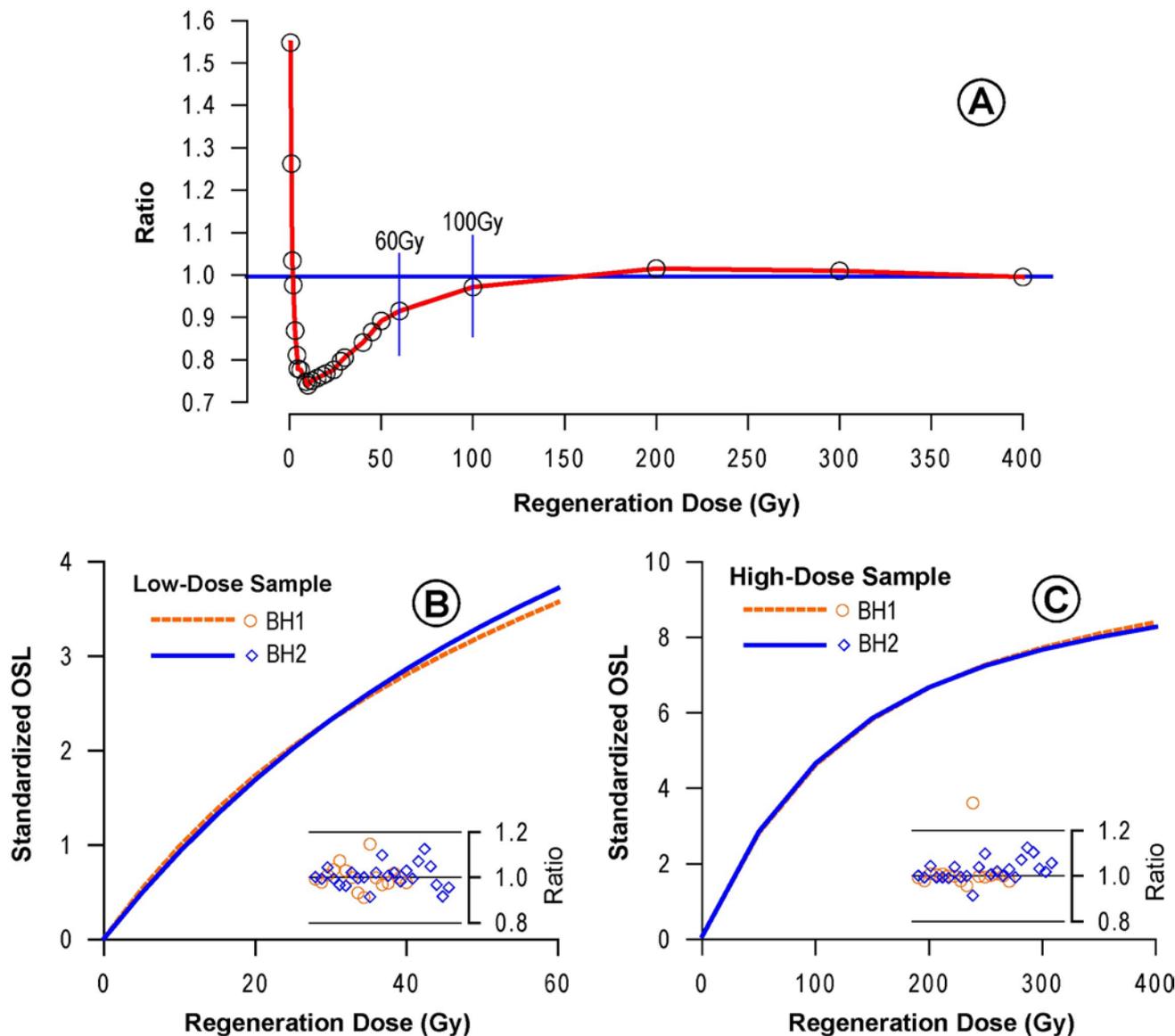


Fig. 6. A, Determination of the boundary for the low- and high-dose samples. The doses were firstly estimated using a combined dataset of all samples, then divided by the original data, and plotted against the original data. B and C are average growth curves of BH1&BH2 samples for low- and high-dose quartz, respectively, estimated from averaging standardized OSL data within samples. Inserts in B and C show ratios between D_e s based SAR and SGC methods: (1) Estimate a SGC based on BH1 dataset, and then estimate D_e s of BH2 dataset, displayed as circles in B and C; (2) Estimate a SGC based on BH2 dataset, and then estimate D_e s of BH1 dataset, displayed as diamonds in B and C. Two equations are displayed in figures, indicating similarity of growth curves between cores. D_e offsets are smaller than 20%, suggesting independence of SGC method on different cores.

come from changes in the sedimentary process but rather from variation in D_e . As shown in **Figs. 2-5**, the ascending trend in D_e was dependent on the regeneration doses, which becomes significant in coarse silty quartz. This has been found in many previous studies (e.g. Lai *et al.*, 2007; Long *et al.*, 2010). According to Yi *et al.* (2012c) and Chen *et al.* (in press) the coarse silty grains are the most dynamics materials in sediment transport, the sedimentary processes of these grains were more complex than others and likely resulted in larger scatter. Because of contamination with heavy minerals (Lai *et al.*, 2007), some grains might have low saturation doses and so create the low- and high-dose samples that share different dose-response curves.

In order to examine the statistical fidelity of a SGC in different dataset, four average growth curves from BH1 and BH2 cores were estimated, and they are shown in **Figs. 6B and C**, respectively. Each curve was then used to estimate the D_e of the other curve. D_e offsets were below 20%, suggesting independence of SGC method on different cores. The growth curves of the BH1 and BH2

cores for low- and high-dose samples are all overlapping, indicating that it is possible to combine these two datasets for use in SGC applications.

Standardized growth curve construction for coarse-silt quartz

Two SGCs were developed by combining data from all twenty-four samples (**Fig. 7**):

1) For low-dose samples:

$$Y = 3.161 \times \left(1 - e^{-\frac{X}{37}}\right) + 0.018 \times X + 0.032 \quad (3.5)$$

The explained variance $R^2=0.999$, the freedom-adjusted explained variance $R^2_{adj}=0.998$, the standard deviation of the regression $s.e.=0.044$, and all coefficients were significant at $p<0.01$.

2) For high-dose samples:

$$Y = 6.940 \times \left(1 - e^{-\frac{X}{111}}\right) + 0.004 \times X + 0.090 \quad (3.6)$$

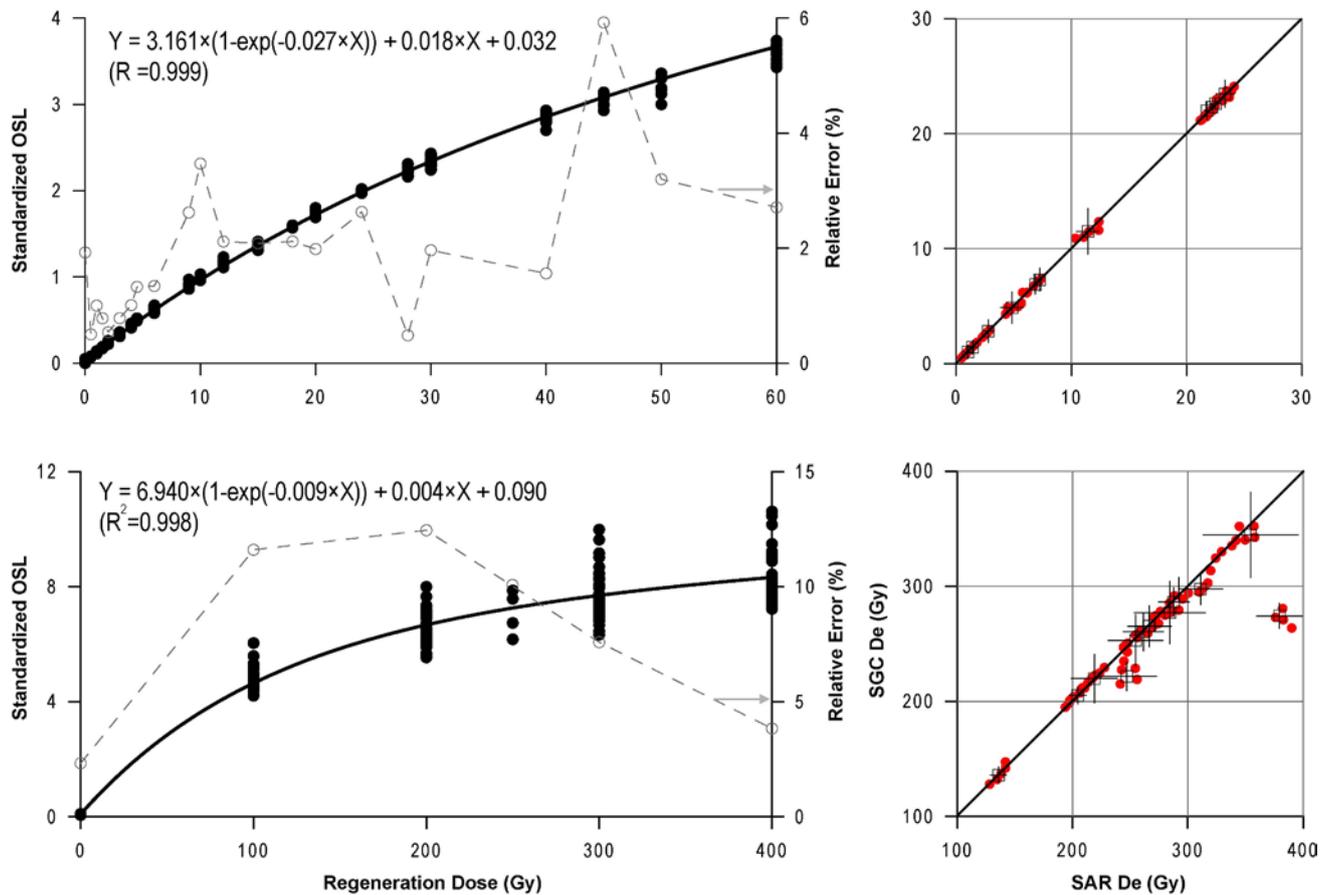


Fig. 7. The SGC constructed using the coarse-silt quartz of twenty-four samples from BH1&BH2 cores (A&C, left panel) for both low-dose (A, upper panel) and high-dose (C, bottom panel). Also shown are the fitting equations of SGC with R^2 values, estimated from averaging standardized OSL data within samples. The recuperation is negligible for these samples, as the fitting line almost goes through the original point. In all cases, the error bar is smaller than the symbol (right-hand y-axis). Comparison of D_e s between SAR and SGC methods (B&D, right panel) for both low-dose (B, upper panel) and high-dose (D, bottom panel). Solid cycles in the right panels indicate individual D_e s.

The explained variance $R^2=0.999$, the freedom-adjusted explained variance $R^2_{adj}=0.997$, the standard deviation of the regression $s.e.=0.161$, and all coefficients were significant at $p<0.01$.

As shown, all regeneration dose points stuck closely to the fitting line. The constants of 0.032 and 0.090 implied that the recuperation for these samples was negligible. In this way, the construction of $SGC_{coarse-silt}$ was found to be proper.

Comparison of D_e s calculated using $SGC_{coarse-silt}$ and SAR

A graphic comparison between D_e s from $SGC_{coarse-silt}$ and SAR is displayed in Fig. 7. All low-dose samples matched each other, but they all showed a much larger variance in D_e s from $SGC_{coarse-silt}$ than from SAR. High-dose samples with doses under 300 Gy matched with each other, but some of the samples with doses over 300 Gy showed marked underestimation of in D_e when $SGC_{coarse-silt}$ was used. Statistical comparison between D_e s from $SGC_{coarse-silt}$ and SAR methods was conducted using One-Way ANOVA. The results were displayed in Table S1. They indicated that only three samples, specifically BH1-15B, BH1-19A and BH2-17B, whose doses were all larger than 300 Gy had significantly different D_e s. This implied that most of D_e s from SGC method were statistically similar to those produced using SAR.

4. CONCLUSIONS

Standardized growth curves have been shown to be reliable in estimation of D_e values if a cautious and thorough approach is taken through the utilization of regional data sets (Telfer *et al.*, 2008). The key points resulting from these experiments with SGCs are as follows: (1) For both the MAR and SAR OSL dating protocols, the SGC method was found to be applicable, providing reliable estimations of D_e . (2) Fine-sand quartz with different palaeodoses had approximately similar dose-response curves, but attention should be paid dating “old” sediments. (3) For coarse-silt quartz, the dose-response curves can be divided into two groups, low-dose (≤ 60 Gy) and high-dose (≥ 100 Gy). Growth curves of low-dose quartz were similar, facilitating the use of SGC in D_e estimations, but errors were large. For high-dose (100-300Gy) quartz, the SGC was also reliable but showed large amounts of uncertainty when D_e was larger than 300Gy. We therefore suggest that SGC could be employed in the dating of coastal sediments with either MAR or SAR OSL protocols and either fine-silt, coarse-silt or fine-sand quartz.

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APPENDIX

The Table S1, the One-Way ANOVA results of estimated D_e s between MAR/SAR and SGC methods from various quartz of three cores drilled from the south Bohai Sea, China, is provided as Supplementary Material and is available in electronic version of this article at <http://dx.doi.org/10.2478/s13386-013-0103-z>.

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